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Concerns about phytoplankton bloom trends in global lakes

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Satellite remote sensing has been widely used to monitor the water quality of inland and coastal environments. Using satellite data from the Landsat 5 Thematic Mapper (L5TM), Ho et al.¹ showed an increase in peak summertime bloom intensity in 68% of the 71 large lakes worldwide from 1982 to 2012. However, we question the veracity of their finding for at least two reasons: (1) satellite-derived reflectance in a single near-infrared (NIR) band is not a reliable proxy for bloom strength, and (2) the infrequent satellite observations from L5TM make it difficult to draw statistically meaningful conclusions.

Ho et al.¹ argued that the L5TM-estimated bloom intensity (B_{NIR}) (see equation 2 in Ho et al.¹), which is basically the reflectance in the NIR band, represents near-surface phytoplankton biomass. However, this argument became questionable when examining the spectral and chlorophyll *a* (Chl*a*; a key indicator for phytoplankton biomass²) datasets collected from 15 lakes in China³ with varying eutrophic status (Chl*a* ranging between 1.5 and 222.6 mg m⁻³; see Extended Data Fig. 1 and Supplementary Information). The complex relationship between spectral reflectance and Chl*a* concentrations were also demonstrated by Spyarakos et al.⁴ using in situ data from around the world. Theoretically, the signal in the NIR band can be attributed to various water constituents in addition to algal blooms, and the contributions from suspended sediments and the presence of aquatic plants could be two of the most common perturbations in inland lakes. Ho et al.¹ attempted to mask out sediment-rich waters with the use of hue but, as detailed later, our analysis suggests that the hue defined in ref.¹ does not accurately represent the colour of a water body.

Bloom strength tends to be substantially overestimated in sediment-rich waters. Examples from two of the lakes studied in Ho et al.¹ (Fig. 1) show that the B_{NIR} value of the high-turbidity, low-algae pixels was higher than that of the algae-present pixels within the same images. The examination of historical images (through both true-colour images and spectral features) shows that L5TM observations have captured sediment plumes in at least 58 (82%) of the 71 studied lakes, and these plumes could be incorrectly labelled as algal blooms owing to their high B_{NIR} (see Extended Data Fig. 2). As supported by previous studies using data from both of the lakes studied in Ho et al.¹ and from other global coastal/inland waters, the NIR reflectance in turbid waters can be substantially enhanced (see Extended Data Table 1). In inland lakes, episodic meteorological (for example, wind and precipitation) and hydrological (for example, riverine discharge) events can strongly influence sediment concentrations⁵, as exemplified by previous studies in Lake Erie⁶ and Lake Okeechobee⁷ in the USA and Hongze Lake⁸ in China (three lakes examined in ref.¹). Therefore, the effect of water turbidity on B_{NIR} should be evaluated carefully.

Similar to high sediment loads, the growth of aquatic vegetation can lead to overestimation of bloom severity. Pixels with high B_{NIR} —in

particular, vegetated waters rather than bloom areas—were also found within the same lakes (see Fig. 1a, b), where massive submerged plants have previously been reported⁹. The reason is that algal blooms and submerged vegetation share similarly high NIR reflectance (see Extended Data Fig. 3). Moreover, previous studies with datasets collected across various global regions and plant species also showed markedly increased NIR reflectance due to the presence of submerged vegetation (see Extended Data Table 2). Indeed, a literature search revealed that of the 71 studied lakes, 41 (58%) contain abundant aquatic plants (see Extended Data Table 3) and their effects on B_{NIR} should have been considered.

A hue-based mask (equations 3 and 4 in Ho et al.¹) was designed to exclude potential contamination from sediments. However, this approach has failed in numerous cases (see Extended Data Fig. 2). This is mainly due to the inclusion of the atmospheric signals in the calculation of hue, that is, the hue was estimated using the top-of-atmosphere reflectance. Thus, this hue represents the colour of the combined signal of the atmosphere and the water. As shown in Extended Data Fig. 4, atmospheric molecular scattering alone could dominate the top-of-atmosphere reflectance for water bodies in the blue band¹⁰. The method (Fmask)¹¹ used to determine lake surface area could lead to substantial underestimations of bloom severity. As the examples in Fig. 1c–e and Extended Data Fig. 5 show, severe bloom areas failed to pass the Fmask and were excluded in further B_{NIR} calculations. Indeed, examination of the lakes studied in ref.¹ showed that most of the bloom scums were missed owing to the improper use of Fmask. This is because intense blooms often cause high normalized difference vegetation index (NDVI) values that can exceed the threshold used by Fmask (for example, NDVI < 0.1) to identify water pixels¹¹.

Furthermore, the interannual dynamics of lacustrine algal blooms are difficult to characterize by infrequent L5TM observations¹², and cloud coverage poses further challenges associated with optical satellite remote sensing. Statistically, the global mean daily cloud-free probability is 33%, with seasonal differences of <5% (ref.¹³). In other words, when L5TM overpasses 23 times in a year because of its 16-day revisit period, the annual mean number of cloud-free observations for a given location is only ~7.5 even without any other unfavourable observational conditions (such as sun glint). As a compromise between data availability and result fidelity, Ho et al.¹ excluded those years with fewer than three valid images in five summer months. We replotted a time series of algal bloom areas in Taihu Lake that was produced by Hu et al.¹² (see Extended Data Fig. 6), which was obtained using daily observations from the Moderate-resolution Imaging Spectroradiometer (MODIS) satellite (revisit period of about one image per day) between 2000 and 2008. Of the >300 cloud-free daily MODIS images within the nine-year period,

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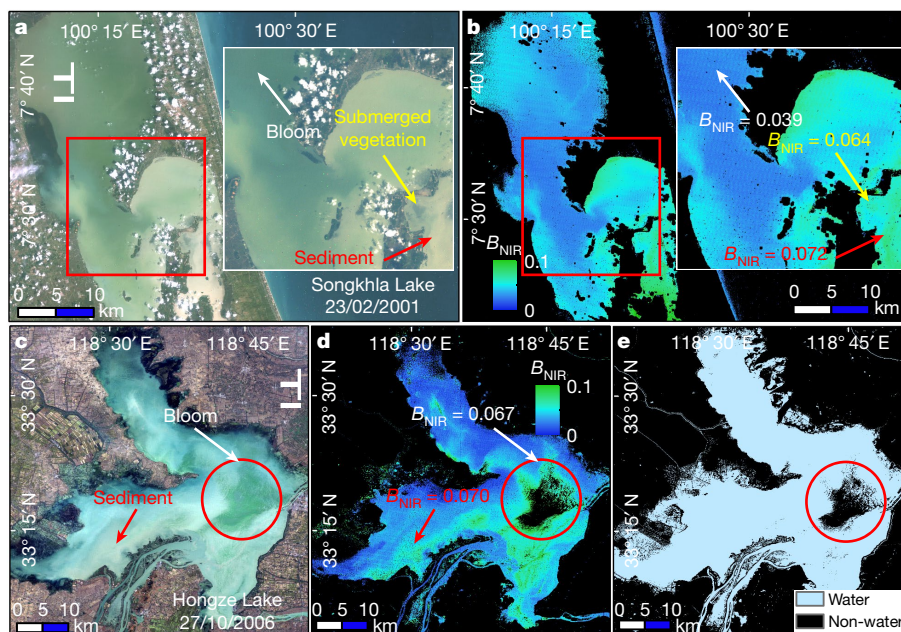


Fig. 1 | Examples showing the problems associated with LSTM-based bloom intensity (B_{NIR}) estimated with equation 2 in Ho et al.¹. a–d, LSTM true-colour composites and corresponding B_{NIR} map for Songkhla Lake in Thailand (a, b) and Hongze Lake in China (c, d). e, Water mask determined by Fmask¹¹ for Hongze Lake using the same image as in c. Areas with either high sediment loads (yellowish in the true-colour images, indicated by red arrows) or the presence of submerged vegetation (dark in the true-colour images, indicated

by yellow arrows) exhibit higher B_{NIR} values than the bloom-occurring pixels (greenish in the true-colour images, indicated by white arrows). An intense bloom in Hongze Lake (within the red circle) was erroneously classified as non-water by Fmask and excluded in the B_{NIR} map (d). The red squares in a, b indicate the inset location. More examples of these problems are shown in Extended Data Figs. 2, 5.

only 24 shared the same overpassing dates as LSTM. Furthermore, detecting a bloom on the basis of remote-sensing imagery depends strongly on wind, given that the fraction of the satellite-observable surface bloom in relation to the total phytoplankton biomass is also a function of wind speed^{14,15}. Owing to the unpredictable nature of cloud occurrence and wind speed, the temporal dynamics of bloom features were difficult to characterize with LSTM datasets.

Our results suggest that the use of LSTM-based B_{NIR} by Ho et al.¹ as a proxy for algal bloom strength is questionable for the majority of the lakes examined in their study. The incorrect use of a water mask algorithm (that is, Fmask) also leads to the omission of the most severe blooms with floating scum. The use of limited Landsat observations (often one cloud-free image every 1–2 months) is problematic for drawing statistically meaningful conclusions. Therefore, the trends in phytoplankton blooms for the 71 global lakes derived by Ho et al.¹ appear unrealistic.

Reporting summary

Further information on research design is available in the Nature Research Reporting Summary linked to this paper.

Data availability

The Landsat data can be obtained from the US Geological Survey at <https://glovis.usg.gov>. The in situ spectral and Chl α data used in this paper are available in the Supplementary Information.

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Competing interests The authors declare no competing interests.

Additional information

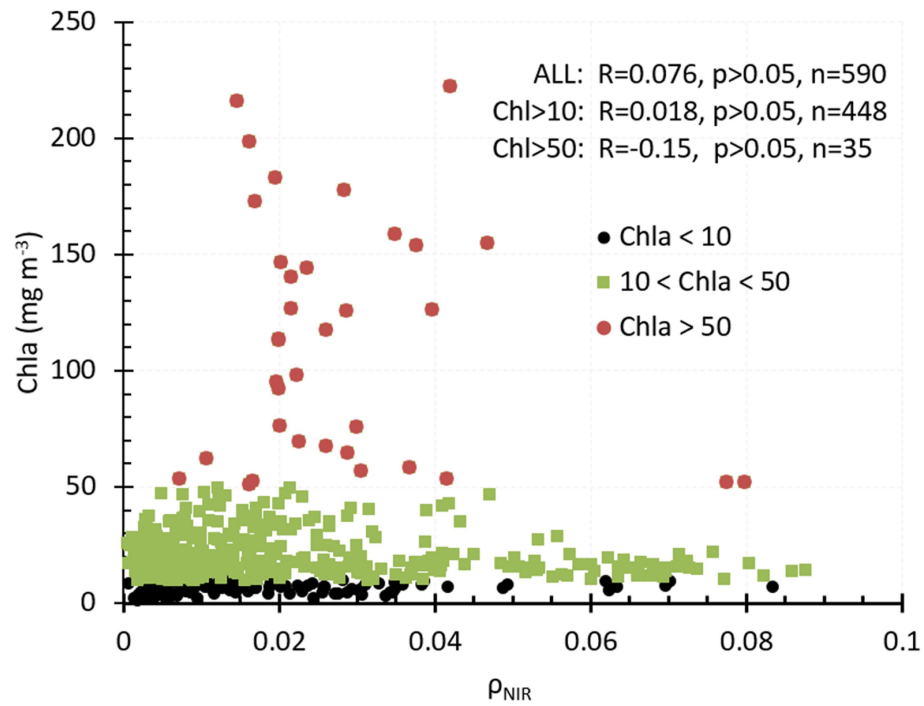
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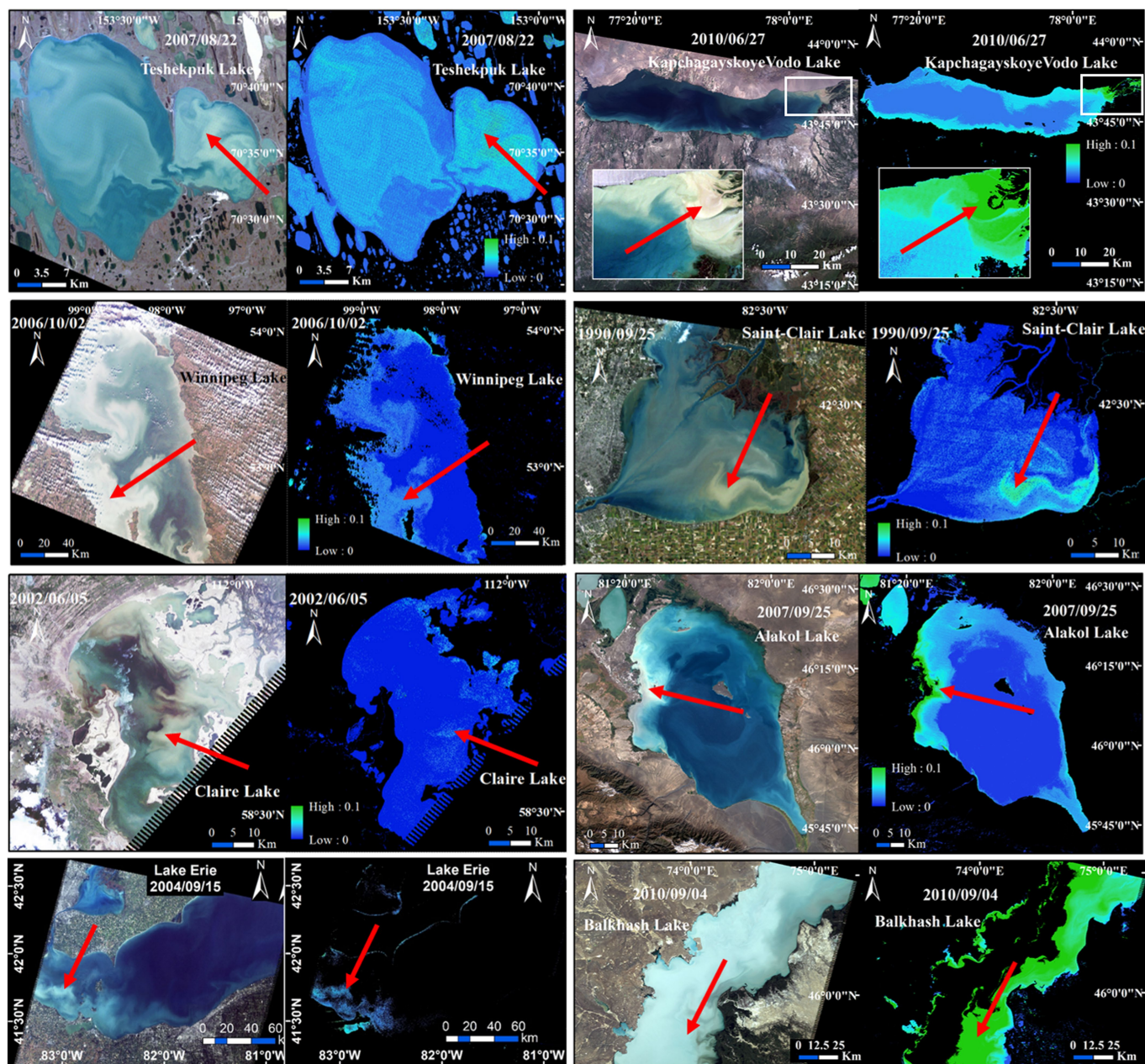
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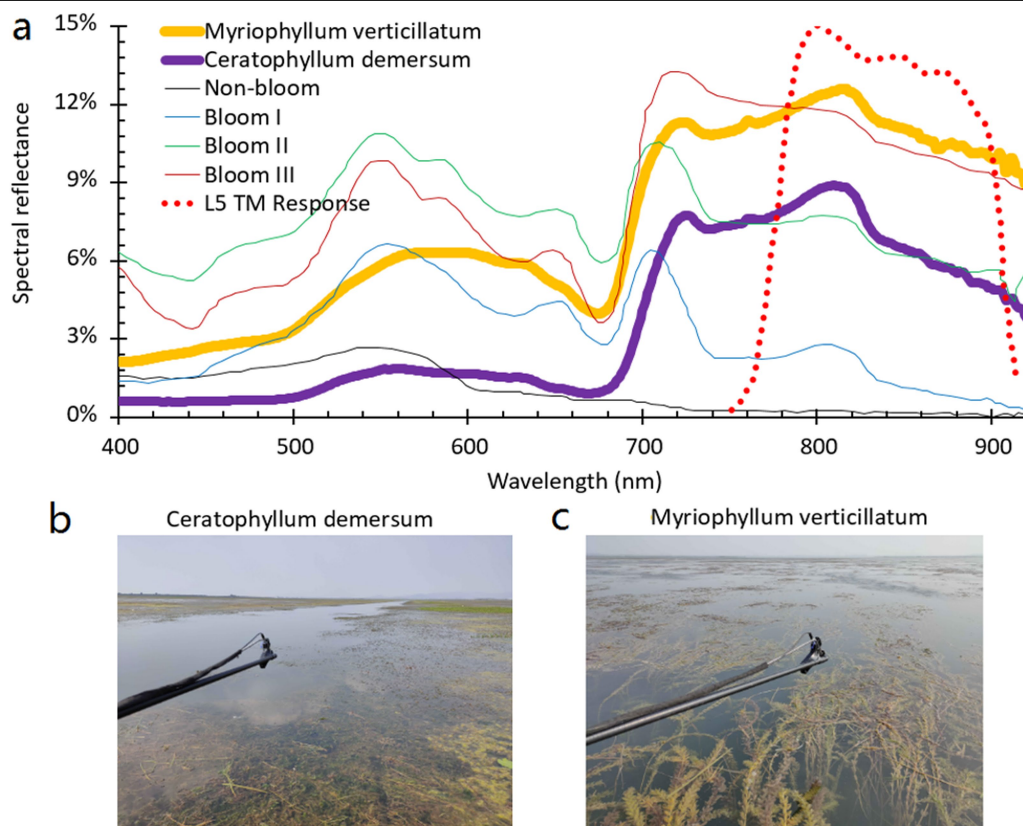
Extended Data Fig. 1 | Relationship between the surface reflectance in the NIR band (ρ_{NIR}) and Chl a. The correlations for different Chl a ranges (colour-coded) and individual lakes are non-significant ($P>0.05$). The data are from in situ measurements collected from 15 lakes in China across waters with

varying eutrophic status. ρ_{NIR} values are the equivalent LSTM NIR reflectances aggregated using in situ hyperspectral measurements and the LSTM spectral response function (see the aggregation method in Kalmen et al.⁹⁶).



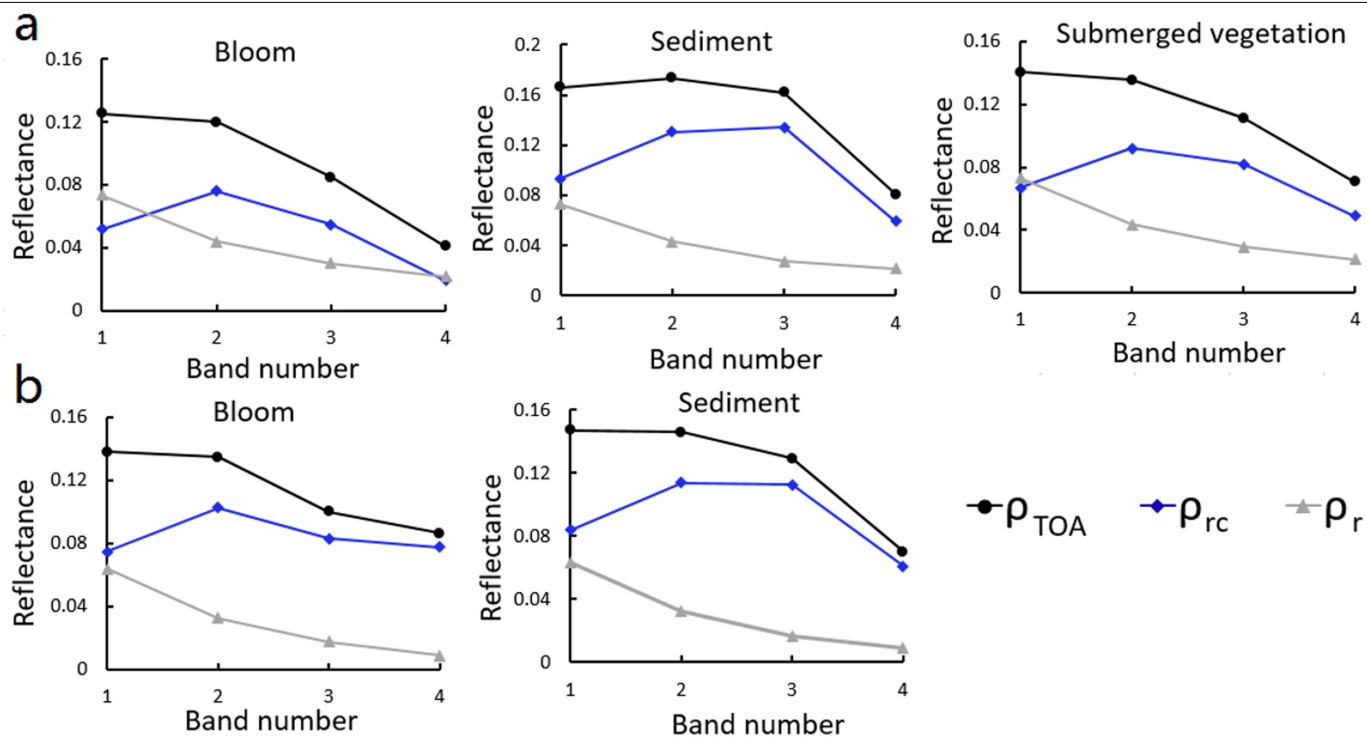
Extended Data Fig. 2 | Examples showing the effects of high sediment loads on the bloom intensity (B_{NIR}) calculations in eight of the lakes studied in Ho et al.¹. The left panels of the paired images show the true-colour composites for LSTM images, and the right panels show the corresponding B_{NIR} maps after applying the hue and Fmask masks. The sediment plumes (indicated by red

arrows) with high B_{NIR} values (~ 0.1) could still be classified as intense blooms with the hue mask defined in Ho et al.¹. The examination of historical LSTM images show that sediment plumes could occur in at least 58 (82%) of the 71 lakes studied in Ho et al.¹.



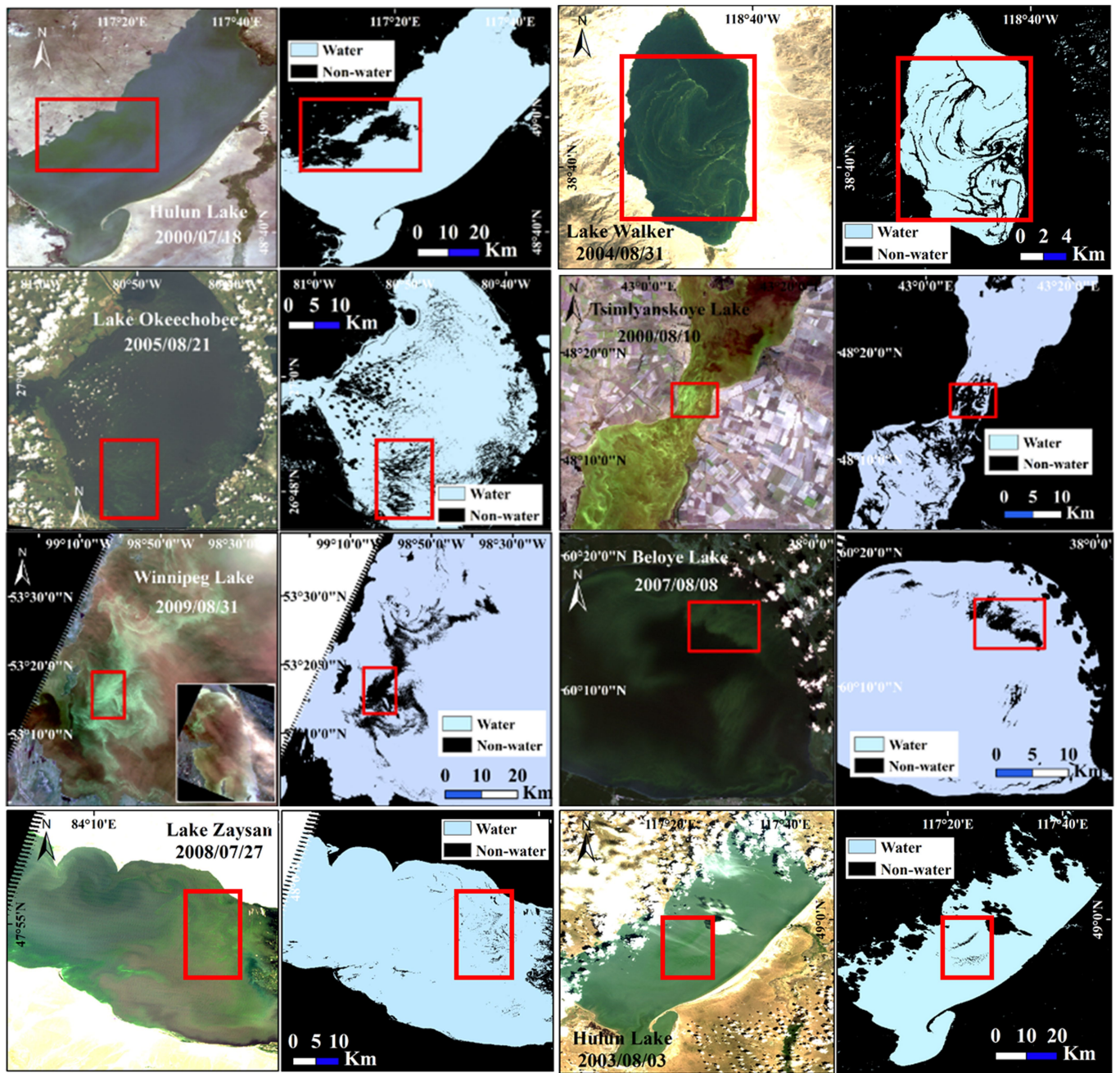
Extended Data Fig. 3 | Reflectance spectra of submerged vegetation.
a, The spectral reflectances of two types (*Ceratophyllum demersum* and *Myriophyllum verticillatum*) of submerged vegetation collected from Taihu Lake in China (a shallow lake ~200 km away from Hongze Lake) on 24 October 2019, using the PSR+3500 field-portable spectrometer manufactured by Spectral Evolution. Also plotted are the spectral reflectances of different

blooms and the normalized spectral responses in the L5TM NIR band, which were obtained from extended data figure 7 in Ho et al.¹. The spectral features of submerged vegetation, particularly the reflectance in the NIR band, are very similar to those of intense phytoplankton blooms. **b**, **c**, Photographs taken while conducting the in situ measurements.



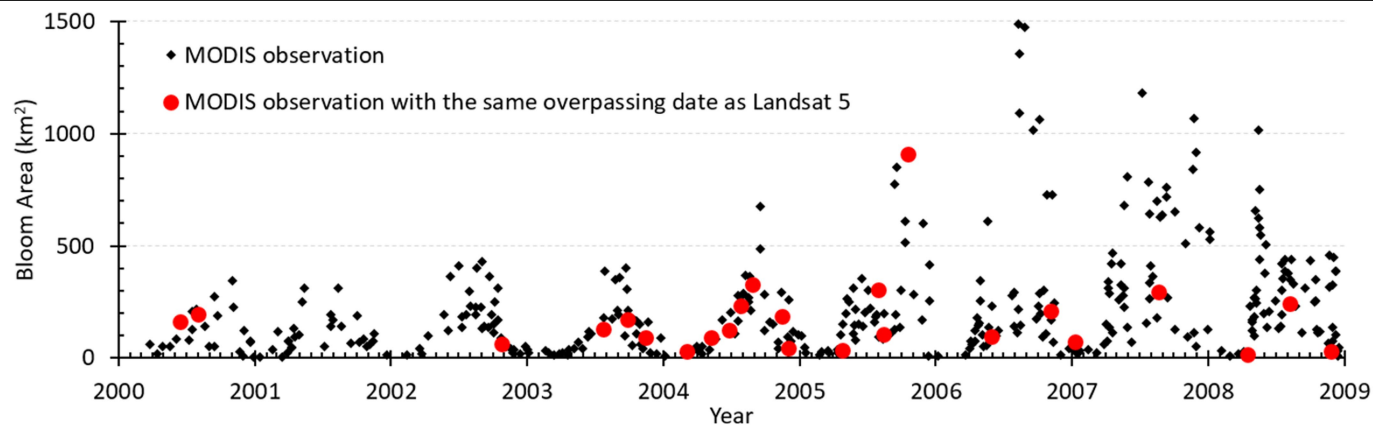
Extended Data Fig. 4 | Spectral features of different types of waters in L5TM images. a, b, The spectral data were obtained from the arrow-indicated pixels in Fig. 1 (a from Songkhla Lake and b from Hongze Lake). ρ_{TOA} is the

top-of-atmosphere reflectance, ρ_r is the reflectance from molecular scattering (or Rayleigh scattering, estimated using the method in Gordon¹⁰) and ρ_{rc} is the difference between ρ_{TOA} and ρ_r .



Extended Data Fig. 5 | Examples of pixels with intense blooms erroneously masked by Fmask in eight of the lakes studied in Ho et al.¹. The left panels of the paired images show the true-colour composites for the LSTM images, and the right panels show the resultant separation of pixels determined using

Fmask. Clearly, intense blooms (greenish in the red squares) have been classified as other classes instead of as water. The examination of the lakes studied in ref.¹ showed that most of the severe blooms with surface scum were missed owing to the improper use of Fmask.



Extended Data Fig. 6 | Daily areas of algal bloom in Taihu Lake between 2000 and 2008, determined using MODIS observations by Hu et al.¹². Red points represent MODIS observations with the same overpassing dates as L5TM (that is, daily MODIS observations have concurrent L5TM images) and

indicate the difficulty in characterizing long-term bloom dynamics. For example, whereas black dots show a clear increase in bloom area after 2005, such a trend is difficult to capture with the red dots.

Extended Data Table 1 | Previous studies with in situ datasets that showed substantial effects of water turbidity (or total suspended sediments, TSS) on the reflectance of the water column in the NIR band

References	Location(s)	Range of turbidity (NTU) /TSS (mg/L)
Büttner et al. ¹⁶	Balaton Lake, Hungary	69.0-591.0 mg/L
Bukata et al. ¹⁷	Saint-Clair Lake, Canada and United States	2.5-20.0 mg/L
Nas et al. ¹⁸	Beysehir Lake, Turkey	0.1-15.5 mg/L
Binding et al. ¹⁹	Lake Erie, United States	0.18-28.26 mg/L
Matthews et al. ²⁰	Zeekoevlei Lake, South Africa	0.03-50 mg/L
Wang et al. ⁷	Lake Okeechobee, United States	1-300 NTU
Kaba et al. ²¹	Tana Lake, Ethiopia	~0-240.0 mg/L
Hamed ²²	Nasser Lake, Egypt	0.75-78.4 NTU
Zeng & Binding ²³	Winnipeg Lake, Canada, and Lake Erie, Canada and United States	0.01-31.6 mg/L
Mikkelsen ²⁴	Four coastal regions, Denmark and Ebro River, Spain	0.5-24.6 mg/L
Dekker et al. ²⁵	Frisian lakes, Netherlands	1.6-255.0 mg/L
Doxaran et al. ²⁶	Gironde Estuary, France	35-2072 mg/L
Koponen et al. ²⁷	Four lakes in Finland	0-30 NTU
Liu et al. ²⁸	Middle Yangtze River, China	23.4–61.2 mg/L
Sterckx et al. ²⁹	Scheldt River, Belgium	13-336 mg/L
Oyama et al. ³⁰	Lake Kasumigaura, Japan	17.6-47.9 mg/L
Tarrant et al. ³¹	Roosevelt and Bartlett Pleasant Lake, United States	0.30–13.4 mg/L
Nechad et al. ³²	Southern North Sea, Europe	1.24-110.27 mg/L
Chen et al. ³³	Apalachicola Bay, United States	1.29–208 mg/L
Knaeps et al. ³⁴	Scheldt, Belgium and the Netherlands	15-402 mg/L
Long & Pavelsky ³⁵	Peace–Athabasca Delta, Canada	0-4000 mg/L
Giardino et al. ³⁶	Lake Maggiore, Italy	0.5-100 mg/L
Feng et al. ³⁷	Yangtze Estuary, China	4.3-1762.1 mg/L
Dorji & Fearn ³⁸	Simulated datasets	0.01-7000.0 mg/L
Dogliotti et al. ³⁹	Southern North Sea; Guyana coastal waters, Scheldt, Gironde, France; and Río de la Plata estuary, South America	1.8-988 mg/L
Han et al. ⁴⁰	European coastal waters; French Guiana; Eastern Vietnam Sea; China Yellow Sea; and northern Canada	0.15-2626.0 mg/L
Yu et al. ⁴¹	Gulf of Mexico and Massachusetts Bay, United States; Yangtze Estuary, China; European coastal waters; the Río de La Plata Estuary, South America	0.2-2068.8 mg/L

The bold text indicates studies of lakes that were also included in Ho et al.¹. We note that this table does not include all related studies, because a complete list would be too long to present here. Data from refs. ⁷¹⁶⁻⁴¹.

Matters arising

Extended Data Table 2 | Previous studies with in situ datasets that showed that NIR reflectance could be substantially enhanced owing to the presence of submerged vegetation

References	Location(s)	Major vegetation species
Zhang et al. ⁴²	Honghu Lake, China	<i>Potamogeton maackianus</i> Benn., <i>Myriophyllum spicatum</i> , <i>Hydrilla verticillata</i> Royle, <i>Ceratophyllum oryzetorum</i> Kom. and <i>Potamogeton lucens</i> Linn.
Vahrmäe et al. ⁴³	Baltic Sea, Europe	<i>Cladophora glomerata</i> , <i>Furcellaria lumbricalis</i> , and <i>Fucus vesiculosus</i>
Dogan et al. ⁴⁴	Lake Mogan, Turkey	<i>Potamogeton pectinatus</i> , <i>Najas</i> sp and <i>Myriophyllum spicatum</i>
Yuan & Zhang ⁴⁵	Chongming Island, China	<i>Myriophyllum spicatum</i>
Yadav et al. ⁴⁶	Lake Biwa, Japan	Unknown
Pu et al. ⁴⁷	Florida coast, United States	<i>Syringodium filiforme</i> , <i>Thalassia</i> <i>testudinum</i> , and <i>Halodule wrightii</i> .
Visser et al. ⁴⁸	River Wylfe and River Frome, UK	<i>Myriophyllum spicatum</i> , <i>Ranunculus fluitans</i> and <i>Potamogeton pectinatus</i>
Watanabe et al. ⁴⁹	Ferreira stream, Brazil	<i>Ceratophyllum demersum</i>
Giardino et al. ⁵⁰	Lake Trasimeno, Italy	<i>Potamogeton pectinatus</i> and <i>Myriophyllum</i> <i>spicatum</i>
Oyama et al. ⁵¹	Lakes Kasumigaura, Inba-numa and Tega-numa, Japan	<i>Trapa natans</i>
Santos et al. ⁵²	Sacramento San Joaquin River Delta, United States	<i>Myriophyllum spicatum</i> and <i>Egeria densa</i>
Luo et al. ⁵³	Lake Taihu, China	<i>Vallisneria spiralis</i> , <i>Ceratophyllum demersum</i> , <i>Potamogeton malaianus</i> , <i>P. maackianus</i> , and <i>Hydrilla verticillata</i>
Hou et al. ⁵⁴	25 lakes, China	Unknown
Brooks et al. ⁵⁵	Lake Huron, United States	<i>M. spicatum</i> and <i>Myriophyllum sibiricum</i>
Fritz et al. ⁵⁶	Lake Starnberg, Germany	<i>Chara</i> spp. and <i>Potamogeton</i> spp.
Ghirardi et al. ⁵⁷	Lake Iseo, Italy	<i>Vallisneria spiralis</i> and <i>Najas marina</i>
Niroumand-Jadidi ⁵⁸	Sarca River, Italy	Unknown
Wilson et al. ⁵⁹	Atlantic coast of Nova Scotia, Canada	<i>Zostera marina</i>

This table does not include all related studies, because a complete list would be too long to present here. Data from refs. ⁴²⁻⁵⁹.

Extended Data Table 3 | List of lakes studied in Ho et al.¹ with abundant submerged vegetation identified

ID	Lake name	Country	References
1	Clear	United States	Niemeier and Hubert ⁶⁰
2	Claire	Canada	Toshner and Region-Brule ⁶¹
3	Simcoe	Canada	Depew et al. ⁶²
4	Bay	Philippines	Vicencio and Buot Jr ⁶³
5	CaboraBassa	Mozambique Zimbabwe	Bond and Roberts ⁶⁴
6	Balaton	Hungary	Istvánovics et al. ⁶⁵
7	Saint-Clair	Canada United States	French III ⁶⁶
8	Dauphin	Canada	Balesic ⁶⁷
9	Khanka	Russia China	Li et al. ⁶⁸
10	Hongze	China	Liu et al. ⁶⁹ ; Shengzhao ⁷⁰
11	Alexandrina	Australia	Ward and Talbot ⁷¹
12	Bosten	China	Wang and Dou ⁷²
13	Okeechobee	United States	Havens et al. ⁷³
14	Poopo	Bolivia	García et al. ⁷⁴
15	Hulun	China	Fang et al. ⁷⁵
16	Songkhla	Thailand	Sompongchaiyakul et al. ⁹
17	Qinghai	China	Chen ⁷⁶
18	Gyaring	China	Pen ⁷⁷
19	Ulungar	China	Li ⁷⁸
20	Ngoring	China	Pen ⁷⁷
21	Se-lin	China	Wang and Dou ⁷²
22	Kariba	Zimbabwe Zambia	Machena ⁷⁹
23	Aral-Sea	Kazakhstan Uzbekistan	Aladin et al. ⁸⁰
24	Winnebago	United States	Gabriel and Bodensteiner ⁸¹
25	Erie	Canada United States	Badzinski et al. ⁸²
26	Baikal	Russia	Chepinoga et al. ⁸³
27	Chilka	India	Jaikumar et al. ⁸⁴
28	Cha-jihNan-mu-tso--Zhari-Namco	China	Wang and Dou ⁷²
29	Beloye	Russia	Krivoronogov et al. ⁸⁵
30	Sasykkol	Kazakhstan	Romanova and Kazangapova ⁸⁶
31	Chapala	Mexico	Villamagna et al. ⁸⁷
32	Balkhash	Kazakhstan	Imentai et al. ⁸⁸
33	Izabal	Guatemala	Barrientos ⁸⁹
34	Urmia	Iran	Tehranchi et al. ⁹⁰
35	Nicaragua	Nicaragua	Davies ⁹¹
36	Alakol	Kazakhstan	Romanova and Kazangapova ⁸⁶
37	Victoria	Tanzania Uganda Kenya	Cheruiyot et al. ⁹²
38	Sevan	Armenia	Heblinski et al. ⁹³
39	Beysehir	Turkey	Beklioglu et al. ⁹⁴
40	Nasser	Egypt Sudan	Green ⁹⁵
41	Edward	Zaire Uganda	Green ⁹⁵

Data from refs. ^{9,60-95}.

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Software and code

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Data collection The satellite data were obtained from the U.S. Geological Survey at <https://glovis.usg.gov>. The in situ spectral and Chla data were collected by authors.

Data analysis ArcGIS (version 10.8) and ENVI (version 5.4) were used to analyze the satellite images

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Study description	This study questioned the veracity of the phytoplankton bloom trends in global lakes derived from Landsat measurements by Ho et al., 2019.
Research sample	Two separate samples were selected. 1) in situ data were used to demonstrate the use of NIR reflectance as a proxy for algae bloom strength is questionable, and 2) satellite images were used to show the erroneous bloom detections using the method presented by Ho et al. 2019.
Sampling strategy	The Landsat data used in our study were from the same lakes that were studied by Ho et al. 2019, and acquisition dates for the data were also fall within their studied summertime.
Data collection	Cruise surveyed in situ data were collected by the authors in 15 large lakes on the Yangtze Plain of China.
Timing and spatial scale	In situ data from 15 lakes in China were collected across four different seasons between 2005 and 2019, and Landsat data covers several lakes that were studied by Ho et al. 2019.
Data exclusions	No data were excluded from analysis.
Reproducibility	Our results could easily be reproduced with existing datasets.
Randomization	The Landsat data were randomly selected from USGS, while only images with small cloud cover were used. All the in situ spectral and Chla data from our filed surveys were used in the study.
Blinding	Not applicable in our study.
Did the study involve field work?	<input type="checkbox"/> Yes <input type="checkbox"/> No

Field work, collection and transport

Field conditions	No precipitation and strong wind occurred when sampling.
Location	15 Lakes on the Yangtze Plain of China, Lat: 29°-30°N, Lon: 112°-122°E
Access & import/export	Not applicable in our study.
Disturbance	NASA-recommended ocean optics protocol (Mobley 1999) was followed when conducting reflectance measurements, to minimize the uncertainties from reflection at the air-water interface.

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<input checked="" type="checkbox"/>	<input type="checkbox"/> MRI-based neuroimaging

Reply to: Concerns about phytoplankton bloom trends in global lakes

<https://doi.org/10.1038/s41586-021-03255-2>
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REPLYING TO L. Feng et al. *Nature* <https://doi.org/10.1038/s41586-021-03254-3> (2021)

In Ho et al.¹, we used three decades of high-resolution Landsat 5 (LSTM) satellite imagery to investigate long-term trends in intense summer-time near-surface phytoplankton blooms for 71 large lakes globally to reveal a global exacerbation of bloom conditions. In the accompanying Comment², Feng et al. question whether the implemented algorithm is “a reliable proxy for bloom strength” owing to the strong effects of suspended sediments and aquatic vegetation, and whether “the infrequent satellite observations from LSTM make it difficult to draw statistically meaningful conclusions”. We agree that these are important considerations, and they have been examined in detail. In Ho et al.³, the algorithm was found to outperform ten other candidate algorithms (see table 1 in ref. ³). The 11 algorithms were evaluated using a total of ten metrics that assessed the algorithms’ ability to capture bloom magnitude, interannual variability and seasonal variability when evaluated against both in situ biovolume observations⁴ and remote-sensing imagery from MERIS and MODIS⁵. In Ho et al.¹, we further evaluated the algorithm by assessing its ability to reproduce previously reported spatial gradients across 48 region pairs in 22 lakes (see section ‘Validation of well known spatial gradients in bloom intensity’ and supplementary table 2 in ref. ¹). In addition, to evaluate whether trends in the abundance of constituents such as suspended sediments or aquatic vegetation could be misinterpreted as trends in bloom intensity, we assessed the algorithm’s ability to capture both the direction and the timing of changes in bloom intensity for seven lakes for which such information was available in the literature (see section ‘Evaluation of bloom intensity time series and trends’ in ref. ¹). In all analyses, the algorithm was shown to meaningfully and accurately represent long-term trends in bloom intensity for lakes for which such information was available. Although we recognize that modern remote-sensing instruments⁶ will be able to track lacustrine algal blooms even more accurately in the future, we maintain that the analysis presented in Ho et al.¹ using LSTM represents the best possible analysis of global phytoplankton bloom intensity over the historical period going back to the 1980s.

In elaborating on the above two topics, Feng et al.² present measurements of near-infrared (NIR) reflectance versus chlorophyll *a* co-mingled across 15 unidentified lakes in China (see extended data figure 1 in ref. ²), whereas in Ho et al.¹ we explicitly state that while the approach implemented therein “has proven effective in identifying the extent of near-surface intense phytoplankton blooms, we emphasize that the retrieval of concentrations of specific bloom severity metrics (for example, chlorophyll *a*) is beyond the scope of this study”. In Ho et al.¹ we also explain that the implemented approach was designed specifically to eliminate the need to compare absolute magnitudes across lakes—a barrier that has limited cross-lake syntheses in the past. For these two reasons, the data presented by Feng et al.² are not directly relevant to the arguments presented in Ho et al.¹.

Feng et al.² then elaborate that “the effect of water turbidity on B_{NIR} should be evaluated carefully”. We agree, and this was indeed done in Ho et al.^{1,3}. As stated above, the algorithm was validated across over two dozen lakes, including an extensive discussion of suspended sediments in Ho et al.³. In figure 1 of ref. ² (and extended data figures 2, 4), Feng et al. use true-colour composite examples from lakes in Ho et al.¹ to argue that sediments can be a confounding factor for the efficacy of the algorithm from Ho et al.¹. However, the simple occurrence of sediment plumes does not imply that long-term trends based on the algorithm are incorrect. Indeed, several of the lakes included in extended data table 1 of Feng et al.² (for example, Lakes Balaton, Winnipeg and Erie) were included in the algorithm evaluation in Ho et al.¹ (see figure 4a and extended data figures 9a, 9c in ref. ¹), and their multi-decadal bloom intensity trends were successfully reproduced by the algorithm. Furthermore, visual interpretation of true-colour imagery alone can be misleading as a method for the delineation of blooms, sediment and submerged vegetation. Indeed, past studies that compared inferences from qualitative observations and remote sensing yielded ambiguous results⁷, which is part of a broader challenge in relating in situ and remote-sensing observations⁸.

Feng et al.² then argue that “the growth of aquatic vegetation can lead to overestimation of bloom severity”. Overall, in optically deep waters in the NIR region, it is very unlikely that the water-leaving signal carries any information pertaining to submerged vegetation, owing to strong absorption by pure water⁹. The key question, however, is not whether submerged vegetation could in some cases affect NIR reflectance (extended data table 2 and extended data figure 3 in ref. ²) or whether some of the examined lakes have submerged vegetation (extended data table 3 in ref. ²), but rather whether trends in the abundance of submerged vegetation could affect the reported trends in bloom intensity. As described above, the algorithm was successfully validated in lakes with well documented trends in bloom intensity, including some of those with submerged vegetation listed in extended data table 3 of ref. ². More broadly, we did not find that the observed trends in bloom intensity tracked documented trends in submerged vegetation for lakes for which such information is available. For example, Havens et al.¹⁰ (see figure 5 therein) report the annual spatial extent of submerged aquatic vegetation for Lake Okeechobee for 2002–2015. The trends in submerged vegetation do not track the observed historical trend in bloom intensity in Ho et al.¹ (see figure 4f therein). There is therefore no indication that trends in submerged aquatic vegetation influenced estimates of long-term trends in bloom intensity.

Feng et al.² then raise concerns about the hue mask. We again refer to the above summary of how the algorithm was validated across over two dozen lakes, including the discussion about suspended sediments and turbidity. We agree that surface reflectance data could be used

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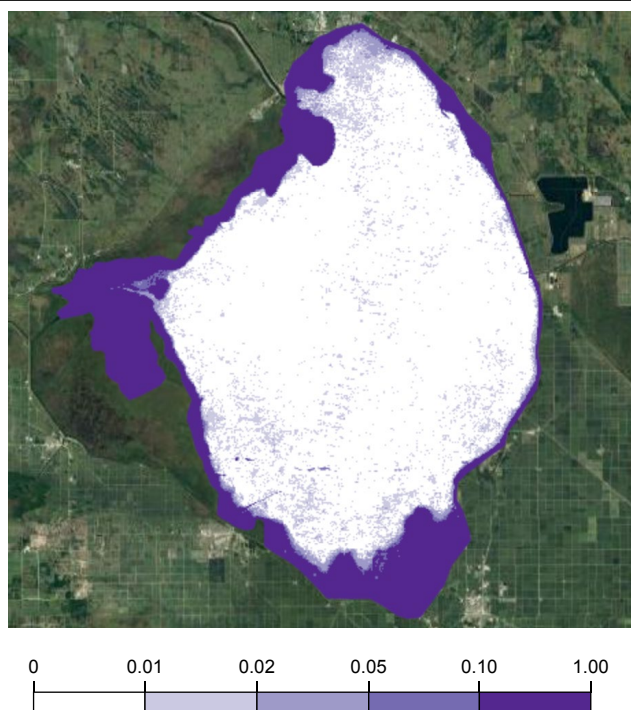


Fig. 1 | Detections of 'land' from Fmask over open water are rare. The map shows the fraction of detections of 'land' from Fmask over Lake Okeechobee from 1984 to 2012 for the months of June to October (242 image composites). Image values are only valid within the Lake polygon boundary. Basemap data from Google¹⁵, TerraMetrics Inc. (www.terrametrics.com).

in future work to further refine the analysis, but such data were not available at the time of the analysis and are not expected to alter the fundamental conclusions of the work.

Next, Feng et al.² comment on the use of the Fmask algorithm for cloud clearing and for identifying water pixels. Extended data figure 5 of ref.² shows examples of cases where the bloom intensity may be so high as to lead the Fmask algorithm to erroneously identify portions of open water as land. Again taking Lake Okeechobee (shown in extended data figure 5 of ref.²) as an example, we examined every scene over the period of record used in Ho et al.¹ and identified the frequency with which pixels in open water were erroneously identified as land using the Fmask algorithm. We confirmed that this can indeed happen in isolated cases, but this is a rare occurrence. Figure 1 shows the fraction of cases, across the 242 scene composites obtained between 1984 and 2012 for the months of June to October, for which a given 30-m pixel within the lake is identified as land. For the open-water portions of the lake, this fraction is below 3%. This illustrates that the Fmask algorithm performs as designed, with variability in the lake boundary and in the near-shore regions being filtered out, whereas the open water where bloom activity occurs is included in almost all composites. Finally, because in Ho et al.¹ we normalize the bloom time series for each lake, the focus is on long-term relative trends in bloom severity, rather than on pixel-by-pixel classification for a particular scene.

Finally, Feng et al.² raise the issue of frequency of observations and cloud cover, and state that because "the global mean daily cloud-free probability is 33%", only one-third of LSTM overpasses are usable. Setting aside the issue that sun glint, one of the "unfavourable observational conditions" resulting in unusable images, as claimed in ref.², is not possible with LSTM data owing to viewing angles below 7.5°, the percentage quoted by Feng et al.² is in no way representative of the data availability for ref.¹. On average, 75% of the area in the LSTM images over the studied lakes over the study period was cloud-free (Fig. 2a). The vast underestimate of cloud-free area in Feng et al.² is likely due

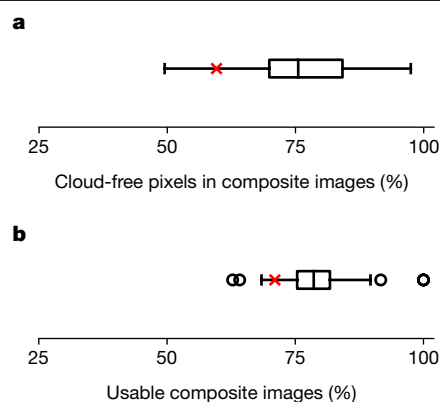


Fig. 2 | The average cloud-free area is large in the LSTM images used in Ho et al.¹ study lakes, and the percentage of usable images is therefore also high. **a**, For each 16-day composite image created for each lake in Ho et al.¹, the cloud-free percentage of the image was calculated as one minus the number of cloud pixels divided by the number of image pixels. Then, for each lake, the cloud-free percentages were averaged across all images to yield an average cloud-free percentage for each lake, presented as a boxplot across the 71 studied lakes. The mean and median across the lakes are 77% and 76% cloud-free pixels, respectively. **b**, For each lake, the percentage of usable composite images was determined by removing individual images with less than 80% cloud-free pixels, or less than the mean observed surface area minus one standard deviation¹. The percentage of usable images is presented as a boxplot across the 71 studied lakes. The mean and median across lakes are both 79% of the composites retained. The red symbols indicate the respective values for Taihu Lake, which was not among the 71 studied lakes. In the boxplots, the boxes indicate the interquartile range (IQR) and the whiskers denote the largest value <1.5×IQR above the third quartile and the smallest value >1.5×IQR below the first quartile. The circles indicate outliers beyond the whisker values.

to the fact that it is based on global cloud cover, which includes high cloud cover over oceans¹¹, does not account for greater cloud cover seasonality over land¹¹ and is calculated from a cloud cover algorithm designed specifically to designate more pixels as cloud versus clear sky in order to be conservative¹².

Extended data figure 6 in ref.² also presents MODIS data for Taihu Lake¹³ (a lake not included in Ho et al.¹) for cloud-free days based on an earlier study¹³, and identify those days with coincident LSTM overpasses to argue that LSTM observations are too sparse to track trends. In Ho et al.¹, we selected lakes on the basis of their inclusion in previous studies that leveraged remote sensing by satellites¹⁴, thus reducing the likelihood that persistent cloudiness obscured the images. Furthermore, we used the Fmask algorithm (https://code.earthengine.google.com/dataset/LANDSAT/LT5_LIT_TOA_FMASK) to objectively identify cloud and cloud shadow pixels (Fig. 2a), whereas the study¹³ referenced in extended data figure 6 of ref.² subjectively "visually examined" images and reported that only images "with minimal cloud cover were chosen and processed"¹³. Also, unlike ref.¹³, we did not exclude all composites with any cloud cover, but only those in which the observed lake surface area was less than 80% of the maximum, or less than the mean minus one standard deviation of the whole surface area of the lake for the time series, whichever was lower. Overall, we were able to use an average of 79% of LSTM images for the lakes included in the study (Fig. 2b). If Lake Taihu had been included in the study, we would have retained 71% of the LSTM images over the study period for that lake (Fig. 2b), or an average of 6.4 images per June–October period, rather than the average 1.4 images per June–October period implied by the red symbols in extended data figure 6 of ref.².

Overall, although we appreciate the careful analysis presented in Feng et al.², none of the issues presented therein had any substantive impact on the analysis and conclusions in Ho et al.¹. Specifically, we saw no evidence that the presence of suspended sediments or aquatic

Matters arising

vegetation affected the long-term trends estimated for the examined lakes, and cloud cover did not reduce data availability to a degree that affected the ability to estimate long-term trends.

Reporting summary

Further information on research design is available in the Nature Research Reporting Summary linked to this paper.

Data availability

Data availability is as outlined in ref. ¹.

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Author contributions J.C.H. and A.M.M. designed the research and analysed the results. J.C.H. and A.M.M. wrote the response with input from N.P. J.C.H. performed the majority of the computations with input from A.M.M.

Competing interests The authors declare no competing interests.

Additional information

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