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Study on the Variation in Evapotranspiration in Different

2 **Period of the Genhe River Basin in China**

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Abstract: Evapotranspiration is an important component and key link of river basin 9 water cycles and plant hydrological processes, and is a core issue in global climate 10 change research. It is not only an important way to understand the energy and water 11 consumption of permafrost regions, but also is an important channel to master the 12 water cycle and energy balance in cold regions. In this paper, multiple linear 13 regression analysis method and weighted comprehensive analysis of major factors 14 method were used to investigate the variation characteristics and impact factors of 15 evapotranspiration in the Genhe River Basin. The results showed the following: (1) 16 The monthly average evapotranspiration in the Genhe River Basin during the 17 freezing-thawing periods in 1980-2017 was 28.29 mm. Compared with the 18 freezing-thawing periods, the total evapotranspiration in the growing seasons was 19 much higher than that in the freezing-thawing periods, with monthly average 20 21 evapotranspiration of 67.71 mm; (2) The main factors affecting evapotranspiration in the Genhe River Basin were precipitation and temperature. During the 22 freezing-thawing periods, the variation in evapotranspiration in May was mainly 23 determined by temperature. In the growing season, precipitation was the main factors 24 affecting evapotranspiration in June. This will lay a foundation for clarifying the 25 relationship between permafrost-climate change-hydrologic cycle in the permafrost 26 active layer during the land surface process, so as to provide some basic data and 27

important scientific basis for the comprehensive study of the hydrologic process and
its impact on climate, ecology, water resources and environment in the permafrost
area.

Key word: Evapotranspiration; Genhe River Basin; Permafrost region; Freezing and
thawing period

33 **1. Introduction**

The evapotranspiration process is a key link in the hydrological cycle that links the 34 atmospheric processes and land-surface processes of the climatic system. In contrast to 35 precipitation, evapotranspiration is a process in which water vapor is transported from 36 37 the surface to the atmosphere. As a core process of the climatic system, evapotranspiration closely links the hydrological cycle, the energy budget and the 38 carbon cycle. Therefore, evapotranspiration studies are important to understand the 39 changes in and effects of the climate and surface (Cammalleri et al., 2010; Vinukollu et 40 41 al., 2011; Li, 2013). At the same time, the variation in and causes of evaporation have very important application value for the assessment of regional basin water resources, 42 crop water requirements, production management, agricultural drought monitoring and 43 ecological environmental problems (such as ecological water demand) (Liu et al., 44 2003). 45

Due to the special climate and environmental conditions in the cold temperate 46 zone of high latitude, the region is less affected by human activities, and the 47 48 eco-hydrological environment is relatively fragile. Evapotranspiration of water, soil and vegetation can reflect the climate change more truly (Li et al., 2019). Calanca et al. 49 50 found that the actual evapotranspiration increased in the high altitude and the area south of the Alps but decreased in the low-altitude area in the northern foreland and the Alps 51 (Calanca et al., 2006). Chen et al. analysed multi-factor combinations that dominated 52 the half-hour evapotranspiration of evergreen coniferous forests across three different 53 climate regions in North America and found that temperature was the most critical to 54

55 the change in evapotranspiration during the growing season (Chen et al., 2018). Findell et al. by developing and applying objective indicators based on physics, found that high 56 evaporation increased the possibility of regional rainfall (Findell et al., 2011). The 57 possibilities of experimentally determining evapotranspiration in the boreal forests of 58 the permafrost zone are limited due to their inaccessibility; therefore, the determination 59 of moisture consumption for evapotranspiration in these regions is performed by 60 computational methods (Budagovsky, 1989; Bondarik et al., 1999; Karpechko et al., 61 62 2010). Previous studies on evapotranspiration mainly included studies on the composition and variation trend of global terrestrial evapotranspiration, the influence 63 mechanism of evapotranspiration and its measurement (Jung et al., 2010; Zhang et al., 64 2016). However, the study area mainly focused on the basins of the warmer regions, 65 such as the middle temperate zone, temperate zone and subtropical zone (Cammalleri et 66 al., 2010). Little attention has been paid by scholars to the variation in and influencing 67 factors of evapotranspiration in the basins of high-latitude and cold regions. In 68 permafrost areas, evapotranspiration of soil, meadow, ice and snow and wetland water 69 70 surface, etc., is an important factor involved in water circulation such as water vapor transport, precipitation, soil infiltration, surface runoff and underground runoff (Li et 71 al., 2019). How meteorological factors affect evapotranspiration in permafrost areas, 72 the variation trend of evapotranspiration in permafrost areas in different periods 73 74 (growing season and freezing-thawing period) and the variation characteristics are all scientific questions to be solved. Due to the existence of complex eco-hydrological 75 processes and fragile ecosystems, alpine regions are extremely vulnerable to damage 76 and are difficult to repair when they are affected by regional climate change and human 77 78 activities. Therefore, it is of great significance to study evapotranspiration in high latitude and cold regions in order to study the response of climate change to water cycle 79 and the eco-hydrological process in permafrost regions. 80

81 The Genhe River Basin is located on the western slope of the northern part of the 82 Greater Xingan Mountains (Fig. 1). It is located in the permafrost region. Most of the

basin is within the temperate, cold and humid forest climatic zone of the Greater 83 Xingan Mountains. The southwestern part of the Genhe River Basin is a temperate and 84 85 semi-humid region with a combination of animal husbandry and agriculture. There are complex ecological and hydrological processes and fragile ecosystems in the Genhe 86 River Basin. In recent years, under the influence of climate warming and human 87 activities (He et al., 2014), ecological environmental changes have complicated 88 evapotranspiration in the basin. From the south to the north, the Genhe River Basin is 89 gradually transformed from medium-temperature grassland to coniferous forest. 90 91 Therefore, considering the distribution of natural ecosystems and the uniqueness of geographical location, it is of great significance to study the regional water cycle in the 92 Genhe River Basin in high-latitude and cold regions. By analysing the variation 93 characteristics of evapotranspiration in the freezing and thawing period and the 94 growing season from 1980 to 2017 in the Genhe River Basin, this paper aims to explore 95 the key factors affecting evapotranspiration and reveal the variation characteristics and 96 97 influence mechanism of evapotranspiration.



98

99

Fig. 1. Location and land use cover of the Genhe River Basin

- 100 2. Methodology
- 101 *2.1 Data*

102 2.1.1 Data Sources

103 The time series of meteorological data are from 1980 to 2017, and the data are from the National Meteorological Information Center (Central Meteorological 104 Administration, 1979) (http://data.cma.cn/). The meteorological data set was obtained 105 from the informatized files of the monthly reports submitted by the Inner Mongolia 106 Autonomous Region, and was compiled based on the relevant regulations of the 107 "Specifications for Surface Meteorological Observations". For details, please refer to 108 the relevant contents of "Specifications for Surface Meteorological Observations". 109 The daily average value was calculated from the data extracted four times (02:00, 110 08:00, 14:00 and 20:00) per day from ground monthly meteorological report data or 111 the real-time database mentioned above. The diurnal data include precipitation, 112 temperature, solar radiation and wind speed. This paper conducts the necessary 113 pre-processing on the data for each meteorological site, including data verification 114 and data interpolation. 115

In this study, the Soil and Water Assessment Tool (SWAT) (Neitsch et al., 2004) 116 117 was used to simulate actual evapotranspiration and soil moisture content in the whole Genhe River Basin. The SWAT model is a semi-distributed basin-scale model that has 118 been used around the world under different conditions (Schuol et al., 2008; Faramazi 119 et al., 2009; Zang et al., 2012). In SWAT, the Genhe River Basin was divided into 29 120 sub-basins and 411 hydrological response units by overlaying elevation, land cover, 121 attributes, 122 soil management, and slope class. In the model, potential evapotranspiration was calculated using Penman formula (Liu et al., 1997). Surface 123

runoff was simulated using a modified SCS curve number (CN) method and snow and 124 melting water calculated by the energy balance equation. Firstly, collect digital 125 126 elevation map of research area, spatial data of land use types, spatial data of soil types, attribute data, meteorological generators, etc. to build a SWAT model database. 127 Various parameters were localized to simulate the monthly runoff from 1980 to 2017, 128 and output sub-watershed output files (SUB), main channel output files (RCH), HRU 129 output files, etc., and then parameter sensitivity analysis and model calibration were 130 carried out. 131

In this study, the measured runoff data of the Genhe River Basin from 1995 to 132 1999 were used for model parameter calibration, and the measured data from 2000 to 133 2009 were used for model parameter verification. SWAT-CUP program was used for 134 parameter sensitivity analysis and parameter calibration of runoff in Genhe River 135 Basin (Abbaspour, 2007). The Nash-Sutcliffe coefficient (Ens) and the coefficient of 136 determination (R^2) were selected to evaluate the goodness of the calibration and 137 validation (Nash et al., 1970). The SUFI-2 method in the SWAT-CUP interface 138 (Abbaspour, 2007) was chosen for parameter optimization. The results of the model 139 simulation verification period (2000–2009) were shown in Fig. 2. R^2 was 0.82 and E_{ns} 140 was 0.79. Therefore, the SWAT model was suitable for the simulation of the 141 hydrological response in the Genhe River Basin. 142



Fig. 2. The SWAT model validation between the observed and simulated
discharge in Genhe river basin from 2000–2009

146 2.1.2 Data Quality Status

143

The quality control codes of the diurnal extreme data and accumulation data of 147 all elements in the data set were provincial quality control codes in the ground 148 monthly meteorological report data files of the same period. Provincial quality control 149 codes were given by the provincial quality control service in the three-level quality 150 control service system of the ground monthly meteorological report data files. The 151 quality control code of the daily average data was taken from the maximum quality 152 control code of the corresponding timing data in the monthly file. The quality control 153 codes of the extreme data and accumulation data were the quality control results of 154 the hourly data files and daily data files of the automatic stations uploaded in real time, 155 which were automatically marked by the quality control software of the stations. The 156

quality control code of the daily average data was taken from the maximum quality control code of the corresponding timing data in the real-time database. Artificial verification was generally carried out on the detected suspected incorrect data, and corrections were carried out on the data that were clearly wrong after verification.

161 *2.2 Method*

162 2.2.1 Data Processing

We used the tessellation polygon method (Burn et al., 2002) to calculate the meteorological data distributions throughout the river basin. We generated tessellation polygons based on the locations of the three weather stations and the boundaries of the Genhe River basin (Fig. 1). We then used the area of each polygon as a weight and calculated the weighted average precipitation throughout the basin.

168 2.2.2 Multiple Linear Regression Analysis

The basic task of the multiple regression equation is to calculate the regression 169 coefficient through a series of dependent variables and independent variables and 170 establish a relational model between the dependent variable and the independent 171 variable. According to a large number of studies (Zhao et al., 2008; Dang et al., 2016; 172 Li et al., 2014), meteorological factors such as solar radiation and rainfall have a great 173 impact on evapotranspiration. The relationship between evapotranspiration and 174 meteorological factors can be expressed by linear regression, and the regression 175 coefficient can directly represent the degree of influence of climatic factors on 176

evapotranspiration (Fan et al., 2019). Therefore, multiple linear regression equations
were adopted in this paper to describe the dependence relationship between the
dependent variable and multiple independent variables. The model was established as
follows:

$$Y_i = \beta_0 + \beta_1 X_{1i} + \beta_2 X_{2i} + \beta_3 X_{3i} + \beta_4 X_{4i} + \beta_5 X_{5i}$$

where Y_i is the evapotranspiration, β_0 is the random error term (constant term), including the influence of other variables not included in the model and random error, β_1 - β_5 is the independent variable regression coefficient, X_{1i} is the precipitation, X_{2i} is the solar radiation, X_{3i} is the temperature, X_{4i} is the wind speed, and X_{5i} is variation in soil moisture content.

187 2.2.3 Weighted Comprehensive Analysis of Major Factors

This paper intends to explore the comprehensive influence degree of major factors on evapotranspiration through the weighted comprehensive analysis of meteorological factors. The calculation steps of this method were as follows (Liu et al., 2020).

192 ①Calculate the correlation between evapotranspiration and key factors such as
193 temperature, solar radiation, wind speed and rainfall and their proportions;

194 ②Carry on the normalized processing to each element value;

195 ③Calculate the comprehensive weighted values of the main influencing factors in196 each period;

197 ④ Calculate the comprehensive influence degree of influence factors on
198 evapotranspiration in each period.

199 The correlation between evapotranspiration and the major factors is expressed by200 the correlation coefficient:

201
$$R_{xy} = \frac{\sum_{i=1}^{n} [(x_i - \bar{x})(y_i - \bar{y})]}{\sqrt{\sum_{i=1}^{n} (x_i - \bar{x})^2} \cdot \sqrt{\sum_{i=1}^{n} (y_i - \bar{y})^2}}$$

where R_{xy} is the coefficient of correlation between x and y; X_i is the evapotranspiration in year i, mm; Y_i is the main factor data of the corresponding year; \overline{X} and \overline{Y} are the mean of evapotranspiration and the mean of meteorological data of the corresponding year, respectively; and i is the year variable.

206 **3. Results**

3.1 Characteristics of Evapotranspiration during the Freezing-thawing Period in the
Genhe River Basin

The monthly average evapotranspiration in the freezing-thawing period in the Genhe River Basin presented a "low-high-low" pattern as a whole, and the variation range was larger than that of the growing season. From April to September, evapotranspiration continued to grow and peaked in September. From September to October, the evapotranspiration decreased continuously and reached a lower value in October.

The monthly average evapotranspiration during the freezing and thawing period was 28.29 mm, and the total evapotranspiration was 113.14 mm, of which 45.84 mm

217	occurred in September, accounting for approximately half of the evapotranspiration
218	during the freezing-thawing periods. Second, the monthly average evapotranspiration
219	in May was 37.48 mm, respectively, accounted for approximately 33% of the
220	freezing-thawing periods. In April and October, the monthly average
221	evapotranspiration, which was 13.91 mm and 15.91 mm, was less (Fig. 3).
222	The monthly evapotranspiration in the Genhe River Basin fluctuated within the
223	range of 5.08–23.76 mm in April, showing an overall fluctuation. The
224	evapotranspiration in the Genhe River Basin in May was distributed in the range of
225	19.84–49.06 mm with high evapotranspiration. The evapotranspiration in the Genhe
226	River Basin in September ranged from 29.93 to 53.74 mm and exhibited fluctuations.
227	The evapotranspiration in the Genhe River Basin decreased significantly in October,
228	fluctuating within the range of 6.5 to 31.8 mm (Table 3).

229

 Table 3. Monthly Average Evapotranspiration in the Freezing-thawing

230

Periods in the Genhe River Basin in 1980-2017

Month	The minimum value	The maximum value	The average value
	(mm)	(mm)	(mm)
April	5.08	23.76	13.91
May	19.84	49.06	37.48
September	29.93	53.74	45.84
October	6.50	31.80	15.91

3.2 Characteristics of Evapotranspiration during the Growing Season in the Genhe
River Basin

233 Compared with the freezing-thawing periods, the monthly average 234 evapotranspiration in the growing season in the Genhe River Basin showed a small 235 change. From June to August, the evapotranspiration increased and reached its peak in 236 August.

The monthly average evapotranspiration in the growing season in the Genhe River Basin was 67.71 mm, and the total evapotranspiration was 203.13 mm. The monthly average evapotranspiration in July and August was 71.59 mm and 77.12 mm, respectively, accounting for more than one-third of the evapotranspiration in the growing season. The monthly average evapotranspiration in June was the least (54.42 mm) (Fig. 4).

The monthly evapotranspiration in the Genhe River Basin fluctuated within the range of 27.7–72.13 mm in June. The monthly average evapotranspiration in the Genhe River Basin increased significantly in July, with a range of 53.62 mm to 81.83 mm. The monthly average evapotranspiration in the Genhe River Basin was within the range of 61.03–95.51 mm in August, with an increase in evapotranspiration compared with June and July (Table 4).

249

 Table 4. Monthly Average Evapotranspiration in the

250 Growing Season in the Genhe River Basin in 1980–2017

Journal Pre-proof									
Month	The minimum value	The maximum value	The average value						
	(mm)	(mm)	(mm)						
June	27.70	72.13	54.42						
July	53.62	81.83	71.59						
August	61.03	95.51	77.12						





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255 Note: The results in Fig. 3 and Fig. 4 were obtained using the monthly average of 1980 to 2017.

3.3 Interannual Variation in Evapotranspiration in the Genhe River Basin in 1980–
257 2017

258	From Fig. 5 and Table A1, the following conclusions can be drawn: from
259	1980 to 2017, the monthly average evapotranspiration showed periodic changes
260	and the average of total evapotranspiration was 317.89 mm. Among them, the
261	highest value of total evapotranspiration appeared in 2014, which was 366.97 mm,
262	and the lowest value appeared in 1987, which was 258.22 mm. At the same time,
263	the five major meteorological factors also showed obvious periodic changes.



264

Fig. 5. Changes in the evapotranspiration and meteorological factors from 1980

266 to 2017

267 3.4 Mechanism Analysis of Evapotranspiration Change in the Genhe River Basin

268 3.4.1 Mechanism Analysis of Evapotranspiration during Freezing-thawing periods in

the Genhe River Basin

The changes in evapotranspiration during the freezing and thawing periods and 270 the growing seasons were mainly affected by energy conditions (such as temperature, 271 solar radiation, water vapor pressure and wind speed) and the water supply conditions 272 of the underlying surface. A large number of studies have shown that the dominant 273 factors for evapotranspiration are different in different regions and seasons. To 274 275 determine the relationship between the evapotranspiration and meteorological factors in the Genhe River Basin, multiple regression coefficients of temperature, 276 precipitation, wind speed, variation in soil moisture content and other related factors 277 and evapotranspiration were calculated during the freezing and thawing periods and 278 the growing seasons (Fig. 3 and Fig. 4). 279

Combining the results of Table 6 and Table 7, we could find that the evapotranspiration changes in the Genhe River Basin during the freezing-thawing periods were mainly affected by three meteorological factors: variation in soil moisture content was negatively correlated with evapotranspiration, and rainfall and temperature were positively correlated with evapotranspiration. The variation in evapotranspiration in April in the Genhe River Basin was mainly caused by solar radiation, which was negatively correlated with evapotranspiration, and temperature,

which was positively correlated with evapotranspiration. The variation in 287 evapotranspiration in May in the Genhe River Basin was mainly caused by 288 temperature, which was positively correlated with evapotranspiration. The change in 289 evapotranspiration in September in the Genhe River Basin was mainly caused by 290 rainfall, which positively correlated with evapotranspiration. The 291 was evapotranspiration changes in the Genhe River Basin during October were mainly 292 affected by two meteorological factors: variation in soil moisture content was 293 negatively correlated with evapotranspiration, and temperature was positively 294 correlated with evapotranspiration. 295

The change in evapotranspiration in the Genhe River Basin was the result of the 296 joint influence of various meteorological elements, and the effects of the 297 meteorological elements on evapotranspiration were different at different scales and 298 periods. Fig. 6 and Fig. 7 show the relationship between evapotranspiration and the 299 weighted values of major factors. As can be seen from Fig. 6, the correlation 300 coefficient of the two was 0.728 in the freezing-thawing period, with a high degree of 301 correlation and good coincidence, indicating that the weighted value of the five 302 influencing factors can well explain the variation in evapotranspiration during the 303 freezing-thawing period. As shown in Fig. 7, in the growing season, the correlation 304 coefficient of the two was 0.352, the correlation was low, and the degree of agreement 305 was general, indicating that the weighted values of the five influencing factors have a 306 weaker interpretation of the changes in evapotranspiration during the growing season. 307

Table 6. Multiple regression equations of major elements and evapotranspiration

309	in the freezing-thawing period in the Genhe River Basin						
		The	The	The	The	The	
Time	Multiple regression equation	significa	significa	significa	significa	significa	
Time		nce of	nce of	nce of	nce of	nce of	\mathbf{R}^2
		X_1	X ₂	X ₃	X_4	X ₅	
Freezing-thawing	Y=6.326X ₁ +22.992X ₂ +1.910	0.000**	0.001**	0.000**	0.008**	0.000**	0.910
period	X ₃ +2.818X ₄ -0.553X ₅ -6.167	0.000	0.001	0.000	0.008		
April	Y=-0.139X ₁ -32.57X ₂ +1.304	0.015	0.000**	0.000**	0.000**	0.442	0.900
Арш	X ₃ +5.227X ₄ -0.059X ₅ +17.323	0.915		0.000	0.000		
May	Y=14.622X ₁ -35.751X ₂ +1.971	0.000**	0.069	0.002*	0.431	0.000**	0 721
Iviay	X ₃ +1.716X ₄ -0.897X ₅ +14.797	0.000	0.009	0.002	0.431	0.000	0.721
September	Y=4.62X ₁ +22.02X ₂ -0.145X ₃	0.013**	0 292	0 807	0 156	0.041	0.464
September	+4.001X ₄ -0.281X ₅ +15.83	0.015	0.272	0.007	0.150	0.041	0.404
October	Y=8.468X ₁ +1.643X ₂ +1.096X ₃	0.007**	0 937	0.002**	0 435	0 000**	0.650
October	$+1.861X_4$ -1.405X ₅ +4.236	0.007	0.751	0.002	0.135	0.000	0.000

310 Note: * indicates significance at the 5% level, and ** indicates significance at the 1% level. Parameter Y is the

311 evapotranspiration; the constant term is the random error term (constant term), including the influence of other

312 variables not included in the model and random error; X_1 is rainfall; X_2 is the solar radiation; X_3 is the temperature;

313 X_4 is the wind speed; and X_5 is variation in soil moisture content.

Table 7. Correlation among evapotranspiration and major factors

Time	X_1	X ₂	X ₃	X_4	X_5
Freezing-thawing period	0.555**	0.232**	0.882^{**}	-0.090	-0.545**
April	-0.547**	-0.736**	0.754**	0.317	-0.550**
May	0.243	0.089	0.567**	-0.133	-0.228
September	0.570^{**}	0.528**	-0.144	0.274	0.215
October	-0.183	-0.212	0.382*	0.314	-0.582**

in the freezing-thawing periods in the Genhe River Basin

316 Note: * indicates significance at the 5% level, and ** indicates significance at the 1% level. Parameter X_1 is

317 rainfall; X_2 is the solar radiation; X_3 is the temperature; X_4 is the wind speed; and X_5 is variation in soil moisture





319

Fig. 6. Relationship between evapotranspiration and the weighted values offactors in the freezing-thawing period in the Genhe River Basin



Fig. 7. Relationship between evapotranspiration and the weighted values offactors in the growing season in the Genhe River Basin

322

325 3.4.2 Analysis of the Mechanism of Evapotranspiration during the Growing season in326 the Genhe River Basin

From the scale of the growing season, the overall increase in evapotranspiration 327 was mainly caused by two meteorological factors: solar radiation and rainfall were 328 329 positively correlated with evapotranspiration. In June, the changes in evapotranspiration in the Genhe River Basin were mainly caused by rainfall, which 330 was positively correlated with evapotranspiration. The change in evapotranspiration in 331 July in the Genhe River Basin was mainly caused by solar radiation, which was 332 positively correlated with evapotranspiration. In August, the changes in 333 evapotranspiration in the Genhe River Basin were mainly caused by rainfall, which 334 335 was positively correlated with evapotranspiration (Fig. 4).

According to the above analysis, the meteorological element positively

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337	correlated with evapotranspiration was mainly temperature. The meteorological
338	elements negatively correlated with evapotranspiration included wind speed and
339	variation in soil moisture content (Table 8 and Table 9).
340	Table 8. Multiple regression equations for major elements in

the evapotranspiration growing season in the Genhe River Basin

341

			The	The	The		
		The				The	
т'		· · c·	significa	significa	significa	• • • •	\mathbf{p}^2
Time	Multiple regression equation	significa	nce of	nce of	nce of	significan	ĸ
		nce of X ₁	nee of	nee or	nee or	ce of X5	
			X ₂	X_3	X_4	•• • • • • • • • •	
Growing	$Y=1.733X_1+101.823X_2+1.553X_3$						
		0.009^{**}	0.000^{**}	0.000^{**}	0.068	0.000^{**}	0.737
season	-4.789X ₄ -0.32X ₅ -29.397						
T	Y=6.8X ₁ +52.709X ₂ +2.122X ₃	0.001**	0.120	0.012*	0.425	0.000**	0 712
June	2 0228 0 4068 22 225	0.001	0.139	0.013	0.425	0.000	0.713
	-2.92374-0.49075-22.233						
	Y=0.874X1+65.51X2+2.933X3						
July	3 1 1 1	0.284	0.015^{*}	0.004**	0.387	0.016^{*}	0.479
-	-3.312X ₄ -0.25X ₅ -28.908						
	$Y=3.63X_1+31.564X+1.751X_3$	<u>ب</u>		4		<i>ч</i>	
August		0.000**	0.181	0.048*	0.060	0.000^{*}	0.706
	$+8.694X_{4}-0.371X_{5}-3.919$						

342 Note: * indicates significance at the 5% level, and ** indicates significance at the 1% level. Parameter Y is the
343 evapotranspiration; the constant term is the random error term (constant term), including the influence of other
344 variables not included in the model and random error; X₁ is rainfall; X₂ is the solar radiation; X₃ is the temperature;
345 X₄ is the wind speed; and X₅ is variation in soil moisture content.

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Table 9. Correlation among evapotranspiration and major factors

Time	X1	X ₂	X ₃	X4	X ₅
Growing season	0.501**	0.748 ^{**}	0.333**	-0.420**	-0.327**
June	0.506**	0.517**	0.110	-0.300	-0.163
July	0.356^{*}	0.365^{*}	0.315	-0.273	-0.250
August	0.576**	0.568^{**}	0.029	0.167	-0.293

in the growing season in the Genhe River Basin

348 Note: * indicates significance at the 5% level, and ** indicates significance at the 1% level. Parameter X_1 is 349 rainfall; X_2 is the solar radiation; X_3 is the temperature; X_4 is the wind speed; and X_5 is variation in soil moisture 350 content.

351 **4. Discussion and Conclusions**

In different climatic zones, the factors affecting forest evapotranspiration vary 352 widely. For example, in the tropical, subtropical and most temperate regions, the main 353 factors affecting forest evapotranspiration are precipitation, temperature and solar 354 radiation (Smith et al., 2013; Cristiano et al., 2015). In cold and temperate regions, in 355 addition to precipitation, temperature and solar radiation, the freezing and thawing of 356 permafrost is also an important factor affecting forest evapotranspiration (Park et al., 357 2008; Tchebakova et al., 2016). The Genhe River Basin has a temperate continental 358 359 climate. Every year in early November, the temperature begins to fall below zero, and the soil begins to freeze. The soil moisture is almost stopped by the supply of rainfall. 360

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The phreatic aquifer in the soil preserved during the growing season moves towards 361 the soil layer, and the soil moisture gradually reaches saturation or supersaturation and 362 363 freezes to form tundra. From the beginning of June, the permafrost begins to melt. The soil moisture content begins to increase gradually, reaching a maximum in July 364 and early August, and permafrost melts to 80% to 90% of the maximum thawing 365 depth. In August, the permafrost continues to melt deep, and plants continue to 366 consume water. Therefore, the soil moisture content begins to decline because the 367 deeper the soil is, the larger the gravel size. The large pores make it difficult for the 368 meltwater to rise to supply the upper layer and at the same time, the precipitation 369 easily infiltrates, so the soil water content begins to decline; the main influencing 370 factors of the change in soil moisture content are permafrost and precipitation (You, 371 2006). From November to March, precipitation (mainly snowfall) is increasing, and 372 the permafrost is thicker. The following spring in April, the temperature during the 373 melt period is high, which is conducive to the rapid melting of the permafrost, and the 374 groundwater level will rise rapidly, which will quickly replenish the water needed for 375 plant germination. The plants turn green, and the transpiration is strengthened. The 376 soil evaporation increases and the evapotranspiration increases. On the other hand, if 377 there is less precipitation in the non-growing season, evapotranspiration is naturally 378 reduced (Wang, 2015). 379

Since the average temperature during the freezing and thawing period is lower than that during the growing season, the lower temperature has a significant inhibitory effect on evapotranspiration, resulting in a lower total evapotranspiration during the

freezing and thawing period; the temperature in summer rises and gradually promotes 383 evapotranspiration (Han et al., 2018; Alexander et al., 2018) so that the 384 385 evapotranspiration is increasing. During the growing season, temperature and rainfall begin to increase from mid-May, the plants begin to grow green, and the transpiration 386 activity strengthens. The soil moisture is also supplemented by rainfall while 387 supplying plant growth and consumption. The increase in the surface temperature 388 causes the increase in soil evapotranspiration at this time and evapotranspiration 389 increases due to the increase in temperature and rainfall. The Genhe River Basin is 390 controlled by cold high pressure in winter, whose temperature is low, and seasonal 391 freezing-thawing occurs in the soil. The soil freezes from the beginning of November 392 every year until the spring in April and May, and the permafrost gradually melts to 393 form the seasonal frozen layer of water, which becomes an important factor for 394 replenishing soil moisture (Wang, 2015). Therefore, during the freezing-thawing 395 periods, evapotranspiration is highly correlated with soil moisture. Because the 396 process of soil freezing is also the process of increasing soil water content in the 397 surface layer (Guo et al., 2002), the soil moisture in the freezing and thawing period is 398 higher than that in the growing season as a whole. During the freezing period, the soil 399 is in a frozen state, the ground temperature is low, and soil evapotranspiration is also 400 rare (Lei et al., 1999). When the permafrost melts, the soil thaws and water infiltrates 401 into the soil, increasing soil moisture. At the same time, evapotranspiration is 402 enhanced under the influence of temperature and other meteorological factors. At this 403 time, due to the correlation between the land surface and atmospheric feedback, the 404

increase in soil moisture will lead to more evapotranspiration (Heerwaarden et al.,
2010). However, the correlation between evapotranspiration and soil moisture in the
growing season is not as high as that in the freezing-thawing period. The situation in
August is particularly prominent. The soil moisture did not change much with the
fluctuation in the evapotranspiration. At this time, other meteorological factors have a
greater impact on evapotranspiration, such as rainfall and temperature.

Research has shown that evaporation depends not only on soil and on meteorological characteristics but also on the composition and age of forest stands (Fedorov, 1981; Fedorov et al., 1990). Therefore, it is necessary to increase studies on the corresponding mechanism of vegetation growth phenology and evapotranspiration and comprehensively consider the meteorological characteristics and forest characteristics to explore the influencing mechanism of forest evapotranspiration in future research.

The selection in this paper of the data from only three meteorological stations to 418 study the variation in evapotranspiration in the Genhe River Basin had a certain 419 impact on the results. Therefore, more data from meteorological stations can be added 420 in future research, which is helpful for understanding hydrological cycle 421 characteristics in high-latitude and cold regions. In this paper, factors such as rainfall, 422 variation in soil moisture content, solar radiation, wind speed and temperature were 423 selected to study the mechanism of evapotranspiration. By analysing the relationship 424 between evapotranspiration and water vapor circulation in cold and temperate regions 425 and the influence of evapotranspiration on freezing-thawing permafrost and the 426

growth of vegetation, this paper can provide the scientific basis and guidance for the 427 continuous dynamic monitoring and estimation of evapotranspiration in high-latitude 428 429 and cold regions. At the same time, this paper provided references for the evaluation of water resources in the alpine region and the study of the impact of climate change 430 (Moses et al., 2018), so as to improve the understanding of the characteristics of 431 evapotranspiration in different periods under the climatic types in this region. It filled 432 in the shortage of research data in high latitude and cold regions' basin. 433 In this paper, the multiple linear regression analysis method and weighted 434 comprehensive analysis of major factors method were used to explore the variation in 435 evapotranspiration in the freezing-thawing periods and the growing season in the 436 Genhe River Basin and to reveal the change mechanism and influencing factors of 437 evapotranspiration in the Genhe River Basin. The conclusions are as follows: 438 Compared with the freezing-thawing periods, the total evapotranspiration in the 1. 439 growing season was higher. During the freezing-thawing periods, the value of 440 evapotranspiration showed irregular fluctuations, which was mainly caused by 441 changes in temperature. Compared with the freezing-thawing periods, the 442 443 evapotranspiration trend in the growing season slowed down, and the range of fluctuation decreased, which was related to the stable level of the temperature 444 and solar radiation in the Genhe River Basin in the growing season. 445 The major elements positively correlated with evapotranspiration was mainly 2. 446

temperature; the major factors negatively correlated with evapotranspirationincluded wind speed and variation in soil moisture content. During the

freezing-thawing periods, the variation in evapotranspiration in May was
mainly determined by temperature. In the growing season, precipitation was
the main factors affecting evapotranspiration in June.

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454 Author statement

Xinqing Lu: Investigation, Data curation, Software, Writing-original draft 455 preparation, Writing—review and editing, Visualization. Chuanfu Zang: 456 Conceptualization, Methodology, Validation, Formal analysis, Investigation, Data 457 curation, Writing-original draft preparation, Writing-review and editing, 458 Supervision, Project administration, Tamara Burenina: Validation, VIsualization, 459 Supervision. 460

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Table A1. Monthly	evapotranspiration	in	the	Genhe	River	Basin	from	1980	to

Year	The minimum value (mm)	The maximum value (mm)	The total amount (mm)
1980	0.00	74.30	259.48
1981	0.00	72.25	304.06
1982	0.00	92.06	337.11
1983	0.00	79.05	311.34
1984	0.00	86.01	339.13
1985	0.00	82.24	327.36
1986	0.00	74.535	301.65
1987	0.00	67.23	258.22
1988	0.01	95.51	357.52
1989	0.00	79.75	328.69
1990	0.00	87.56	358.67
1991	0.00	86.47	337.12
1992	0.00	76.79	310.77
1993	0.00	77.38	331.46
1994	0.00	73.76	320.58
1995	0.00	68.03	311.87
1996	0.00	90.94	335.58
1997	0.00	82.18	320.53
1998	0.00	82.64	331.00
1999	0.00	82.64	335.26
2000	0.00	75.07	308.08
2001	0.00	72.51	270.60
2002	0.00	72.40	295.08
2003	0.00	74.57	295.02
2004	0.00	69.04	280.71
2005	0.00	79.992	330.94

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2006	0.00	75.22	310.85				
2007	0.00	66.78	280.27				
2008	0.00	80.16	306.72				
2009	0.00	76.91	334.70				
2010	0.00	72.81	311.34				
2011	0.00	82.50	325.84				
2012	0.00	79.08	332.97				
2013	0.00	87.74	362.70				
2014	0.00	84.58	366.97				
2015	0.00	83.37	330.58				
2016	0.00	79.06	313.31				
2017	0.00	74.82	305.85				

Highlights

- 1. The evapotranspiration variability in permafrost region have been analyzed in growing season and freezing and thawing period.
- 2. The influencing factors of evapotranspiration in a long term have been explored.
- 3. This study enriches the data of eco-hydrology research in data-deficient areas.

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Declaration of interests

 \boxtimes The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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